AMINO ACID RACEMISATION DATING OF A RAISED GRAVEL BEACH DEPOSIT, SELLICKS BEACH, SOUTH AUSTRALIA

by C.V. Murray-Wallace* & R.P. Bourman*

Summary

MURRAY-WALLACE, C.V. & BOURMAN, R.P. (2002). Amino acid racemisation dating of a raised gravel beach deposit, Sellicks Beach, South Australia. *Trans. R. Soc. S. Aust.* **126** (1), 21-28, 31 May, 2002.

The extent of racemisation (total acid hydrolysate) of the amino acids aspartic acid, glutamic acid, leucine, phenylalanine and valine indicates a minimum age of last interglacial for lossil molluses occurring within a raised gravel beach deposit at Sellicks Beach, South Australia. The base of the raised gravel beach occurs up to 5.5 m above Australian Height Datum (AHD) and possibly indicates 3 m of local uplift since the last interglacial maximum (c, 125 ka; Oxygen Isotope Substage 5e). Emergence of the gravel beach is attributed to ongoing neotectonic uplift of Fleuricu Peninsula.

KLY WORDS: amino acid racemisation, last interglacial, neotectonics, sea-levels, South Australia.

Introduction

A resurgence of interest in recent years in Quaternary emergent shoreline successions has arisen from the increasing ability to determine the age of these features due to technological advances in geochronology (Rutter & Catto 1995; Noller et al. 2000). Similarly, an increasing awareness that coastal successions, particularly those deposited during the last interglaciation (c.125 ka), are sufficiently old to quantify even modest rates of neotectonism, has bolstered this research endeavour. Accordingly, the elevation of last interglacial coastal deposits has been widely used as a benehmark to delineate recent teetonic behaviour at continental scales (Murray-Wallace & Belperio 1991; Ota 1994; Bourman et al. 1999; Zazo et al. 1999). In this work, the age of a raised beach deposit at southern Sellicks Beach, South Australia, is determined based on the extent of racemisation of several amino acids within molluses from the fossil assemblage. In addition, the neotectonic significance of this deposit and its relation to other emergent shoreline deposits on Fleurieu Peninsula is examined.

Materials and Methods

Field investigations

The elevation and lateral extent of the gravel beach deposit was surveyed to Australian Height Datum (AHD) using an automatic level. In addition to a general field description of the deposit, shell samples

were collected for amino acid racemisation dating and to document the fossil molluse assemblage. Species identification followed that set out in Ludbrook (1984).

Amino acid racentisation analyses

Samples of fossil molluses for amino acid racemisation analyses (total acid hydrolysate) were collected from the gravel beach deposit. Shells were removed from the matrix of the deposit and their depth of burial recorded. Analyses were undertaken on specimens of *Patella (Scutellastra) laticostata* Blainville, *Thais orbita* (Gmelin), *Sydaphera unclulata* (Sowerby), *Nerita (Melanerita) atramentosa* Reeve and *Ostrea* sp. Linnacus.

Sediment adhering to the surfaces of shell samples and diagenetically modified aragonite, particularly chalky surfaces, were removed with a dental drill, followed by successive washes in distilled water using an ultrasonic bath. A dilute aeid etch (2 mol HCl) was subsequently undertaken to remove the outer surfaces (c. 10-15% by mass) of the shells that had been in contact with the host sediment. Samples were subsequently hydrolysed for 22 hours at 110° C in 8 mol HCl. Following cation exchange isolation of the amino acid residues, samples were freeze dried and derivatized. Chromatography of the N-penta-fluropropionyl D, L-amino acid 2-propyl esters was performed using a Hewlett-Paekard 5890A Series II gas chromatograph with a llame ionisation detector and a 25 m eoiled, fused siliea eapillary eolumn eoated with the stationary phase Chirasil-L-Val. Full details of the analytical techniques followed in this work are reported elsewhere (Murray-Wallace 1993). Enantiomeric ratios were determined for the amino acids aspartic acid (ASP), glutamic acid (GLU), leucine (LEU), phenylalanine (PHE) and valine (VAL).

^{*} School of Geosciences, University of Wollongong, NSW 2522.

[†] School of Environment and Recreation Management, University of South Australia, Mawson Lakes, SA 5095.

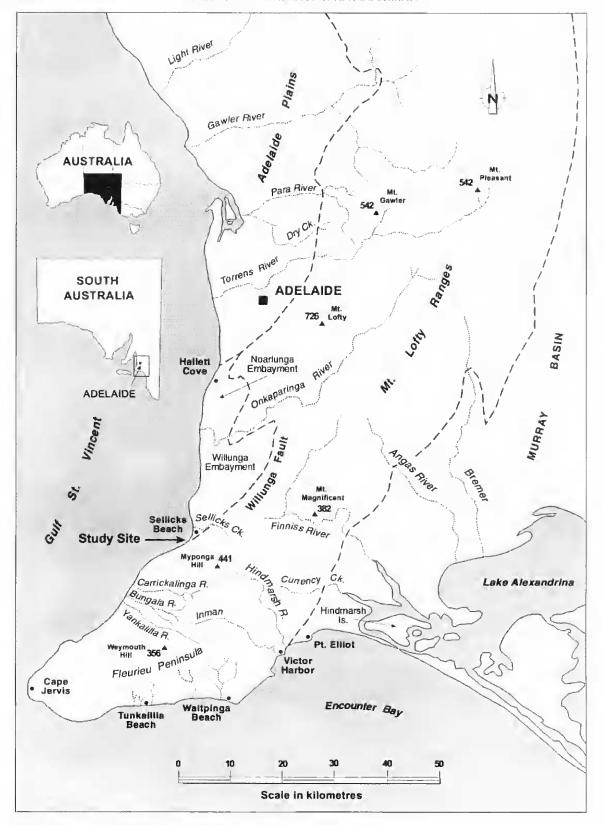


Fig. 1 Location of the raised gravel beach deposit, Sellicks Beach, South Australia.



Fig 2. View Tooking south along southern Sellieks Beach towards the southern Adelaide Hills and coastal eliffs developed on Pleistocene alluvial fair successions. The Jocation of the raised beach deposit, which occurs in the scarp foot zone is indicated by an arrow.



Fig. 4. A shore-normal view of the raised gravel beach deposit at Sellicks Beach. The steeply dipping Ochre Coye llormation is visible near the survey stall. The stall_ which is fully extended, is 5 m long.

Geomorphological Setting And Site Description

The raised gravel beach deposit is situated near the Willunga Fault at the southern-most part of Sellicks Beach (35° 21' 09.8" S; 138° 26' 07.5" E), landward of a modern, gently seaward sloping intertidal shore platform (Figs 1, 2). The modern intertidal platform is approximately 20 m wide in a shore-normal transect, is partially covered with boulders and cobbles and represents a modern analogue for the reliet platform (Fig. 3). An accumulation of boulders and cobbles occurs at the foot of the modern cliff and represents a further modern analogue of the raised beach deposit. The emergent gravel beach facies rests on a strongly croded remnant of a shore platform that is developed on the steeply dipping Oligo-Miocene Port Willunga Beds (Daily et al.



Fig. 3. View looking southwest towards the raised gravel heach deposit at Sellicks Beach. The gravels rest unconformably on the Oligo-Miocene Port Willunga Beds and the Middle Pleistocene Ochre Cove Formation. The gravel deposit dips gently seawards. The unconformity surface represents a relici intertidal shore platform. The modern intertidal platform occurs in the foreground, dips gently seaward and is partially covered by boulders and cobbles. The maximum difference in elevation between the two platforms is 5 m as determined in the most landward exposure, not visible in this photograph. Small, isolated, sea stacks representing crosional remnants of the formerly more extensive Pleistocene shore platform occur within this area (e.g., "a" in the middle distance).

1976), and in part, a steeply dipping portion of the Middle Pleistocene Ochre Cove Formation (Ward 1966; Pillans & Bournan 1996; Figs 3, 4, 5). The gravel deposit occurs within a former scarp foot zone excavated at a time of higher sea level, and abuts fanglomerates of the Ochre Cove Formation (May & Bournan 1984).

The bedrock surface on which the gravel beach facies rests, grades in a seaward direction from 5.55 m to 4.95 m above Australian Height Datum (AHD). The platform extends out seaward from the deposit some 1-1.5 m forming a well-defined bench (Fig. 6). The gravel beach facies crops out over a shore-parallel distance of approximately 50 m, and ranges in thickness between 1 and 1.5 m (Fig. 6).

The gravel deposit is poorly sorted and comprises sub-rounded to subangular clasts of siltstone, quartzite and bryozoal limestone that range from boulder to pebble size, although the modal clast size is boulder-cobble (700-70 mm). The lithoclasts are tightly packed. Numerous entire and fragmental fossil molluses occur within the granular matrix of the gravel deposit.

A pale grey, clean, free-flowing sand is thinly draped over the gravel deposit and the underlying fanglomerates and extends up to 2.5 m above the upper bounding surface of the deposit. The sand also



Fig. 5. Detail from figure 4 showing the rightly packed arrangement of lithoclasts.



Fig. 6. View looking eastnortheast showing part of the shore-parallel lateral extent of the raised beach deposit. The letter "n" denotes the general level of the gravel deposit which is approximately 4 in above the gravel covered footshipe of the small cliff in the foreground. A planated surface representing remnants of an interndal shore platform is visible on the seaward side of the deposit. Pleistocene fanglomerates are evident in the upper right-hand portion of the photograph "b". The raised beach deposit is overlain by a flun veneer of sand which also partially covers the fanglomerate, but is difficult to discern in this photograph.

occurs within the uppermost part of the matrix of the gravel bed near the contact between the gravel and the overlying sand. A thermoluminescence age of 34.0±2.9 ka (W2317) was previously reported for this sandy unit (Bourman et al. 1999). In addition, a radiocarbon age (minimum age) of >30 ka (GaK-6095) has previously been reported for molluses from the gravel beach deposit (May & Bourman 1984).

Results and Discussion

Mollase assemblage

The gravel unit contains a relatively diverse

assemblage of fossil molluses, principally gastropods, within the sediment matrix. Molluses include Potella (Scatellastra) laticostata Blainville, Mactra rufescens (Lamarck), Ostrea angust Sowerby, Monodonta (Austrocochlea) constricta Lamarck, Nerita (Melanerita) atramentosa Reeve. Cymatiella Jesuguri Iredale; Comus sp. Linnaeus, Diloma (Chloraditoma) adelaidea (Philippi). Bembreium melanostuma (Gmelin), Sydaphera undulata (Sowerby) and opercula of Turbo sp., Limnaeus. Many of the shells also occur as large fragments, highly abraded and of unrecognizable affinity. Collectively, the fossil assemblage indicates deposition in an environment comparable to the modern coast at southern Sellicks Beach, with molluses found in sand or attached to rocks, in a relatively sheltered setting of the lower littoral zone (Ludbrook 1984).

Daning

A generally high degree of racemisation (expressed as a D/L ratio) is evident for the five different enantiomeric amino acids measured in each of the fossil molluses from the gravel beach deposit (Table 1). The relative extent of racemisation for the different amino acids, within the single molluse samples, generally follows the relation VAL<LEU<GLU<PHISASP. Similar trends are reported for fossil molluses from United States Pacific coastal plain sites (Lajoie et al, 1980).

Three specimens of Patella (Scatellastra). lancostata (samples UWGA-695, 696 and 763) reveal good concordance in measured chantiomeric ratios (i.e. between-shell D/L ratio variation) with coefficients of variation less than 5.6% for all amino acids for the combined data (actual values include VAL 2.2%, LEU 5.6%, ASP 0.3%, PHE 1.9% and GLU 5%). The consistently lower degree of racemisation for all amino acids in the specimen of Patella sp. (sample UWGA-697), compared with the other three Patella samples is possibly due to the diffusive loss of the more highly racemised, lower molecular weight peptide fraction from the shell carbonate matrix. Accordingly, the degree of racemisation as determined in the total acid hydrolysate, would be disproportionately weighted towards the less racemised, higher molecular weight peptide residues that remain within the shell aragonile matrix. This explanation is consistent with the poorly preserved nature of some of the molluses within the gravel deposit (e.g. chalky appearance).

The high extent of racemisation measured in all the fossil molluses from the raised beach deposit far exceeds values typically determined in Holocene fossils (Murray-Wallace & Bourman 1990; Murray-Wallace & Goode 1995; Murray-Wallace 2000; Table 1). The extent of racemisation in the molluses

TABLE A. Extent of amono acid racemisation (total acid hydrolysate) in fossil molluses from a raised gravel beach deposit. Sellicks Beach and other localities for comparison

Species & Location	Lab, Code			Amino acid D/L ratio ^r		
	or reference	VAL.	LEU	ASP	PHE	GLU
Sellicks Beach,						
raised beach deposit						
Thals arbita (columella)	UWGA-733	0.284 ± 0.021	0.369	-0.556 ± 0.015	+0	۵.
Sydaphera midulata	UWGA-736	-	0.333	0.322 ± 0.036		*
Patella (Sentellastra)	UWGA-697	0.309 ± 0.008	0.370+0.018	-0.613 ± 0.033	0.541 - 0.006	0,549 0,051
laticostata						
Patella (Scutellastra) laticostata	UWGA-696	0.41.2	0.582) 0.021	0.800±0.004	0,798±0.029	0,606 0,002
Patella (Sentellastra) laticostata	UWGA-695	0.405+0.003	0.551+0.007	0,79910,007	0.777±0.017	0.557±0.000
Patella (Scutellastra) laticustata	UWGA-763	0.423±0.008	0.520±0.003	0.80410.001	0.770±0.609	0.611±0.009
Nerita (Melanerita) atramentusa	LIWGA-766	0.38n±0.008	0.411±0.00e	0.702±0.029	0,599±0.005	0.671±0.006
Ostrea sp.	UWGA-768	0.36520.010	0.403±0.013	0.835±0.023	0.727±0.008	0.789±0.026
Late Pleistocene, Glanville Formation, Normanville, SA						
Mactra australis	Bourman <i>et al.</i> (1999)	0,283 (0,011	0.273±0.012	0.590±0.010	40-	0.333±0.006
Hindmarsh Island, SA Mactra australis		0.26±0.003	0,37±0,002	(0.56±0.00]		0,3610,002
Port Wakefield, SA Anadava trapezta	Murray-Wallace	0.32±0.06	0.51+0.02	0.54+0.03		0.43±0.01
Katelysia rhytiphora	(1988)	0.32+0.04	0.51+0.07	0.46±0.02	-	0,38+0,04
Holocene Three Rivers Creek, King Island, TAS						
Patella laticostata (790±60 yr BP; SUA-2927)	Murray-Wallace & Goede (1995)	0.01	0.05±0.02	-	0.04=0.001	0.05±0.001
Sir Richard Peninsula, S						1.
Donay deltoides (22601140 yr BP; SUA-2881)	Murray-Wallace & Bourman (1990)	0.07±0.01	-	0.27±0.01	0.19±0.01	0.12+0.005

Famino acids: VA1 - valine; LEU - leacine; ASP - aspartic acid, PHE - phenylalamne and GLU - glutamic acid.

from the Sellicks Beach deposit also exceeds that apparent for representative examples from the Late Pleistocene Glanville Formation at Normanville and Hindmarsh Island, two localities with comparable current mean annual air temperatures and, as a corollary, two deposits likely to have experienced similar diagenetic temperature histories to the Sellicks Beach deposit, given the caveat that the shells from each deposit remained buried at depths ≥1 m for much of their diagenetic histories (Murray-Wallace et al. 1988; Bourman et al. 1999; Table 1). The Glanville Formation, as originally defined in the

Adetaide region (Ludbrook 1976; Cann 1978) has been correlated with the last interglacial maximum (125 ka: Oxygen Isotope Substage 5e) based on thermoluminescence, amino acid racemisation and uranium-series dating of correlative deposits from other parts of the South Australian coastline (Belperio et al. 1984; Schwebel 1984; Huntley et al. 1993, 1994; Murray-Wallace 2000).

Although the fossil molluses from the raised beach deposit at Sellicks Beach were obtained from near-surface contexts (<50 cm), the geomorphological and stratigraphical evidence suggest that for part of

their diagenetic history, the fossils were more deeply buried (i.e. at least 1 m). However, these molluses will have experienced a higher integrated diagenetic temperature than for fossils that have remained in more deeply buried contexts (Table 1). Current mean annual air temperatures (CMAT) at Sellieks Beach, Normanville and Hindmarsh Island are all approximately 16° C, and 17° C for Port Wakefield.

The extent of racemisation for the majority of amino acids is significantly higher in the molluses from Sellieks Beach compared with those from Normanville and Hindmarsh Island (Table 1). The difference in extent of racemisation is less pronounced when compared with the molluses from Port Wakefield which have experienced a higher diagenetic temperature (Table 1). As current mean annual temperature at the Port Wakefield site is approximately To C warmer than at Sellicks Beach. and given that rates of racemisation are known to increase by up to 20 per cent for such a temperature difference (McCoy 1987), the implication is that the shallow burial depth of the shells at Sellicks Beach has contributed to the high degree of racemisation of amino acids within these fossils.

Amino acid D/L ratios for the molluses from the Sellicks Beach deposit range from the envelope of values representative of last interglacial age to potentially the penultimate interglacial (e. 220 ka; Oxygen Isotope Stage 7), as revealed in a plot of the extent of facemisation against current mean annual temperature (and as a corollary, latitude) (Fig. 7). The lack of clustering and chronological consistency of the data suggests a diagenetic basis for the observed variation in chantiomeric ratios rather than a genuine age variation between shells. The range in D/L ratios for the shells from Sellicks Beach exceeds that typically found for a single isotopic stage (Murray-Wallace 2000).

Although racemisation rates are known to be genus-specific (Miller & Brigham-Grette 1989) this is unlikely to account solely for the higher degree of racemisation in the fossil Patella (Scutellastra) Inticustata from the Sellicks Beach deposit, compared with other genera from the Glanville Furnation. It is therefore concluded that the higher degree of racemisation in the molluses from Sellicks Beach is due to faster rates of racemisation, due to their shallow burial depth during late diagenesis resulting from the progressive exhumation of the deposit, and a genus-effect on racemisation.

As the shells have been subjected to variable burial depths during late diagenesis, the integrated rare expression for racemisation was rearranged with temperature as the subject to assess whether it is possible to induce the high extent of racemisation at ambient diagenetic temperatures over the course of the Hologene. As the amino acid analyses reported

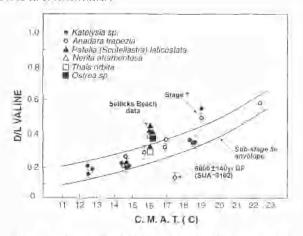


Fig. 7. The extent of valine recemisation front avid hydrolysate) in fossil molluses of last interglacial age (Osygen Isotope Substage Ser from southern Australia plotted against current mean annual air temperature (° C) to illustrate the Sellicks Beach data within a broader. regional context. Details of samples from elsewhere in southern Australia are reported by Murray-Wallace & Belperio (1991) and Murray-Wallace et al. (1999). The amino acid data for the last interglacial molluses are inaccord with the exponential trend of increasing extent of meemisation with higher diagenetic temperatures, and as a corollary, higher gurrent mean annual temperatures. The fossil mollises from the raised beach deposit at Sellicks Beach reveal a broad range in extent of neconisation. from the envelope of the last interglacial to values consistent with a penultimate interglacial age (Oxygen Isotope Stage 7 e. 220 ka). Amino acid results for Halocene and Stage 7 molluses are presented as a framework for comparison.

here were undertaken on different fossils from those used for the radiocarbon assay (-30 ka age; May & Bourman 1984), the integrated diagenetie temperature was calculated to examine the possibility that the radiocarbon age was the result of chance sampling of reworked Pleistocene shells within a Holocene deposit. A minimum age of 7000 years was selected for the calculation, representing the timing of the culmination of the post-glacial marine transgression in southern Australia (Belpeno et al. 2002), and, therefore, the oldest age likely for an undisturbed Holocene coastal deposit. The rationale for this is that the early Holocene is the only time in the Late Quaternary, apart from the last interglacial maximum, that sea level was sufficiently high potentially to form the raised beach. Present sealevel is not sufficiently high to form the deposit. Furthermore, interstadial sea levels of the Late Pleistoeene (Chappell et al. 1996) were significantly below present sea level and would imply rates of tectonic uplift of a magnitude inconsistent with the well-established tectonic framework for the region (Bourman et al. 1999; Belperio et al. 2002).

An average diagonetic temperature required to induce the degree of racconisation measured in the fossils assuming an age of 7 ka was determined thus:

$$T = \frac{5939}{15.77 - \log_{10} \ln \left[\frac{(1+D/L)}{(1-D/L)f} - \ln \left[\frac{(1+D/L)}{(1-D/L)fo} \right] - \frac{1}{2} \frac{1}{(1-D/L)fo} \right]}$$

where T is the absolute temperature (°K), D/L, and D/L, are the enantiomeric ratios of the fossils and their modern equivalents respectively, t is an assumed age (i.e. 7000 years) and 15.77 and 5939 are constants derived from the empirical rate constant expression (Webmitler 1982, 1993). Accordingly, an average diagenetic temperature of 24° C would be necessary to induce the extent of racemisation measured in the three specimens of Patella (Soutellastra) laticostata (UWGA-695, -696 and -763) from the raised beach deposit if they were only 7 ka. A diagenetic temperature of this value is unlikely, however, given that the current mean annual air temperature at Sellieks Beach is approximately 16° C. A prolonged, higher mean annual temperature by as much as 8° C is unlikely over the course of the Holocene for this region (Chappell 1991) Thus, the extent of racemisation measured in the fossil molluses from the raised beach deposit could not have been attained during the Holocene. A penultimate interglacial age is also not favoured, as the gravel beach deposit is unlikely to have survived erosional processes of the last two glacial cycles. On this basis a last interglacial age is favoured for the raised gravel beach deposit at Sellicks Beach:

Neotectonics

The raised gravel beach deposit at Sellicks Beach provides a further opportunity to examine the neotectonic behaviour of Fleuricu Peninsula, Previous investigations have revealed that the region has experienced geologically recent uplift as indicated by the elevation of last interglacial coastal deposits (Bourman et al. 1999).

Although many gravel beach deposits represent relational sea-level indicators (i.e. always form above tidal datum) and are therefore of only modest reliability (Chappell 1987), several attributes of the deposit at Sellicks Beach render it more reliable for quantifying rates of neorectonism. The adjacent

modern intertidal platform has clearly formed within a narrow range of tidal datum and represents an analogous feature to the Pleistocene equivalent. The upper reaches of the modern shore platform are covered by boulders and cobbles presumably accumulated during storm events. However, the steep backing slope of the cliff prevents boulders or finer clasts from being deposited at any significantly higher elevation above tidal datum.

Estimates of a glacio-custatic sea level for the last interglacial (Oxygen Isotope Substage 5c) from Eyre Peninsula suggest a value of 2 m AHD, and represents a particularly reliable datum given the relative tectonic stability of the Gawler Craton upon which much of the Eyre Peninsula coastline has developed (Murray-Wallace & Belperio 1991). Thus, uplift of the Sellicks Beach deposit by as much as 3 m is indicated based on the elevation of the contact between the gravel deposit and the underlying erosional surface of the relict shore platform.

The amount of uplift since the last interglacial maximum, inferred from the deposit at Sellicks Beach (c: 3 m) is less than that observed at Normanville (c. 10 m) to the south of the Willunga Fault (Bourman et al. 1999). The uplift is attributed to the combined effects of pagoing tectonic uplift of the Adelaide Hills and crosional unloading and associated crustal isostatic compensation. Further research is required to model these processes geophysically.

Conclusions

The extent of raccinisation for several amino acids in fossil molluses from a raised gravel beach deposit at Sellicks Beach, South Australia, indicates that the deposit is of Late Pleistocene age, and most likely formed during the last interglacial maximum (c. 125 ka; Oxygen Isotope Substage 5c). The deposit indicates up to 3 m of uplift has occurred in this region since the last interglacial and suggests that the region is still undergoing neotectonic uplift.

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References

BLUDGROU, A. P. HARVEY, N. & BOURMAN, R. P. (2002) Spatial and temporal variability in the Holocene palaeo sea-level record around the South Australian coastline. Sed Geol. (in press).

SMILL, B. W., POLACH, H. A., NETGURER. C. A., Dr Master, D. J., Prescott, J. R., Harls, J. R. & GOSTIN, V. A. (1984) Chronological studies of the Quaternary marine sediments of northern Spencer Gulf.

South Australia. Mar. Geol. 61, 265-296,

BOURMAN, R. P., HELPERIO, A. P., MURRAY-WALLACE, C. V. & CANN, J. H. (1999) A last interglacial embayment fill at Normanville, South Australia, and its neotectonic implications. Trans. R. Soc. S. Aust. 123, 1-15,

CANN, J. II. (1978) An exposed reference section for the Glanville Formation. Quarterly Geological Notes.

Geological Survey of South Australia 65, 2-4.

CHAPPELL, J. (1987) Late Quaternary sea-level changes in the Australian region pp. 296-331 In Tooley, M. J. & Shennan, J. (Eds) "Sea-Level Changes" (Oxford, Basil Blackwell).

(1991) Late Quaternary environmental changes in eastern and central Australia, and their climatic

interpretation. Qual. Sci. Rev. 10, 377-390.

OMURA, A., ESAT, L., McCl. LOGGI, M.,
PANDOLLI, J., OLA, Y. & PILLANS, B. (1996)
Reconciliation of late Quaternary sea levels derived from coral terraces at Huon Peninsula with deep sea oxygen isotope records. Earth Planet. Sci. Lett. 141, 227-236.

Daily, B., Firman, J. B., Fordis, B. G. & Lindsay, J. M. (1976) Geology pp. 5-42 In Twidale, C. R., Tyler, M. J. & Webb, B. P. (Eds.) "Natural History of the Adelaide Region" (Royal Society of South Australia, Adelaide).

HUNTLEY, D. J., HUTTON, J. T. & PRESCOTT, J. R. (1993) The stranded beach-dune sequence of south-oast South Australia: a test of thermoluminescence dating, 0-800 kg. Qual Sci. Rev. 12, 1-20.

(1994) Further de thermoluminescence dates from the dune sequence in the sombeast of South Australia. Ibid. 13, 201-207.

LAIDIE, K. R., WEHMILLER, J. P. & KENNEDY, G. L. (1980) Inter- and intra-generic trends in apparent racemization kinetics of animo acids in Quaternary mollusks pp. 305-340 In Hare, P. E., Hoering, L. C. & King, K. (Lds) "Biogeochemistry of amino acids" (Wiley, New York).

LEDBROOK, N. H. (1976) The Glanville Formation of Port Adelaide, Quarterly Geological Notes, Geological

Survey of South Australia 57, 4-7

(1984) Quaternary mofluses of South Australia, Department of Mines and Energy, South Australia, Handbook No. 9.

MAY, R. M. & BOURMAN, R. P. (1984) Coastal landslumping in Pleistocene sediments at Sellicks Beach, South

Australia, Trans. R. Sov. S. Aust. 108, 85-94. McCox, W. D. (1987) The precision of amino acid geoehronology and paleothermometry. Duat. Sci. Rev. 6, 43-

MILLER, G. H. & BRIGHAM-GREETE, J. (1989) Amino acid geochronology: resolution and precision in carbonate fossils, Quar. Internat. 1, 111-128.

MURRAY-WALLACE, C. V. (1993) A review of the application of the amino acid racemisation reaction to archaeological diffing, The director 16, 19-26.

(2000) Quaternary coastal aminostratigraphy: Australian data in a global context pp. 279-300 In Goodfriend, G. A., Collins, M. J., Fogel, M. L., Macko, S. A. & Wehmiller, J. F. (Eds) "Perspectives in amino acid and protein geochemistry" (Oxford University Press, New York).

& BULDERIY, A. P. (1991) The last interglacial shoreline in Australia - A review. Quar. Sci. Rev. 10, 441-

461.

& BOURMAN, R. J. (1990) Direct radiocarbon valibration for ammo neid racemization dating. Just. J. Farth Sci. 37_365-367.

BELIERIO, A. P., BOURMAN, B. P., CANN, J. H. &C Proces, D. M. (1999) Facies architecture of a Jast interglacial harrier: a model for Quaternary Farrier development from the Coorong to Mount Gambier Coastal Plain, southeastern Australia, Mar. Gool, 158, 177-195.

& Grupt, A. (1995) Aminostratigraphy and electron spin resonance duting of Quaternary coastal neotectonism in Tasmania and the Bass Strait islands. dust J. Earth Sci. 42, 51-67

KIMBER, R. W. L., BELPERIO, A. P. & GOSTIN, V. A. (1988) Aminostratigraphy of the last interglacial insouthern Australia. Search 19, 33-36.

NOLUB, J. S., SOWERS, J. M. & LITTES, W. R. (2000) (Eds) *Quaternary Geochronology: methods and applications*
(American Geophysical Union, Reference Shelf 4, Washington, DC.)

Ora, Y. (1994) Last Interglacial Shorelines in the Western.

Pacific Rim. J. Geog. 103, 809-827.

PHILASS, B. J. & BOLRMAN, R. P. (1996). The Brunbes/Matuyania polarity transition as a chronostratigraphic marker in Australian regoluli studies ACSO J. Anst. Geol. & Geophys. 16, 289-294.

RUTTER, N. W & CATTO, N. R. (1995) (Eds) "Dating methods for Quaternary deposits" (Cicological

Association of Canada, GEOText 2).

Schweißer, D. A. (1984) Quaternary stratigraphy and sealevel variations in the southeast of South Australia pp. 291-311 In Thom, B. G. (Ed.) "Coastal geomorphotopy in Australia" (Academic Press, Sydney).

WARD, W. T. (1966) Geology, geomorphology and soils of the south-western part of County Adelaide, South

Australia, CSIRO Soils Publication No. 23.

WEINBLER, J. F. (1982) A review of amino acid racefurzation studies in Quaternary mollusks: stratigraphic and chronologic applications in coastal and interglacial sites, Pacific and Atlantic Coasts, United States, United Kingdom, Baffin Island and tropical islands. Onat. Ser. Rev. 1, 83-120.

(1993) Applications of organic geochemstry for Quaternary research: aminostratteraphy and aminochronology pp. 755-783 In Engel, M. & Macko, S.

(10ls) "Organic geochemistry" (Plenum, New York). ZAZO, C., SILVA, P. G., GOY, J. L., HILLARI, MARCIA, C., GUALEB, B., LARIO, J., BARDAN, T. & ROZÁLIZ, A. (1999) Coastal uplift in continental collision plate boundaries: data from the Last Interglacial marine terraces of the Gibraltar Strait area (south Spain), Tectonophys. 301, 95